

Improved Initialization of the L-A System? A Need to Revisit the Modeling of Surface States and Fluxes

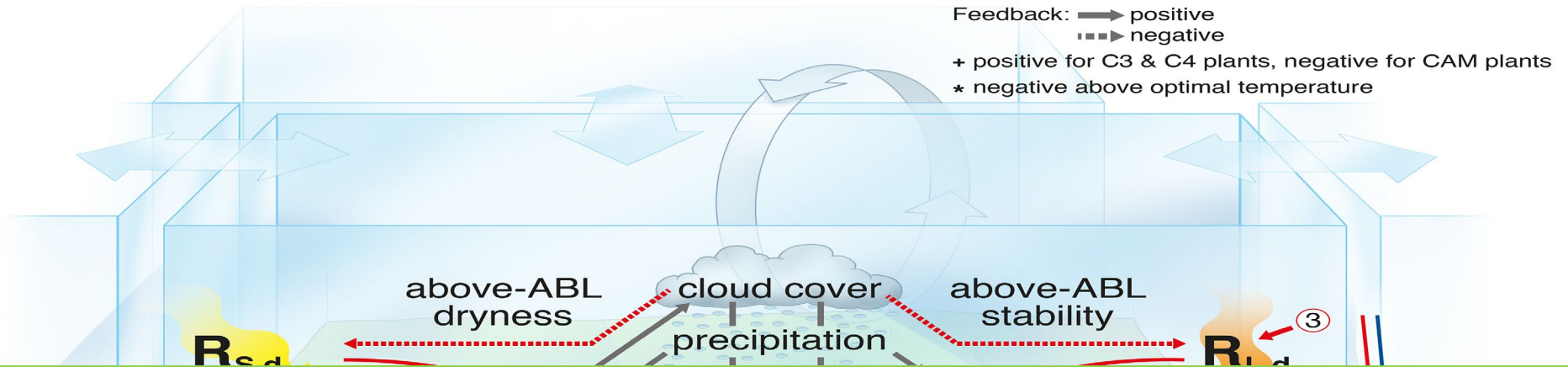
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Volker Wulfmeyer (WGNE representative for GLASS), Nathaniel Chaney and Anne Verhoef (GLASS co-chairs)

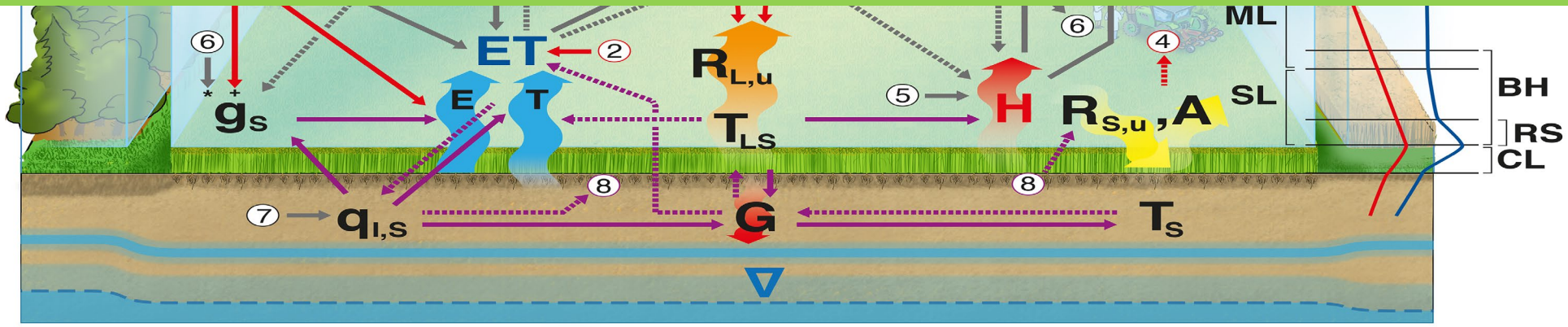
With materials from the GLASS Panel Project Leaders

The Land-Atmosphere System

2



How well does the complexity of land surface modeling lead to significant improvements in the simulation of surface fluxes and land-atmosphere interactions?



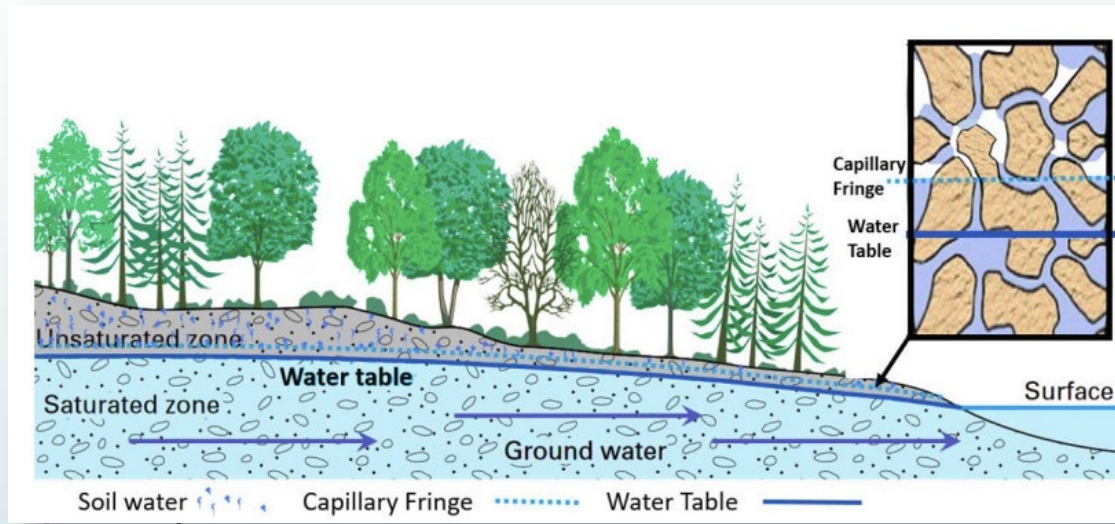
Implications for Increased Land Surface Model Complexity (and Coupled Model Demand)

1. More process representation: Do we see improved modeling of surface states and fluxes?
2. Large increase in the number of model parameters: Is this a good thing?
3. Moving towards high-resolutions (km-scale): Growing challenges in how to initialize these models (spin-up time and data availability/readiness)
4. ...

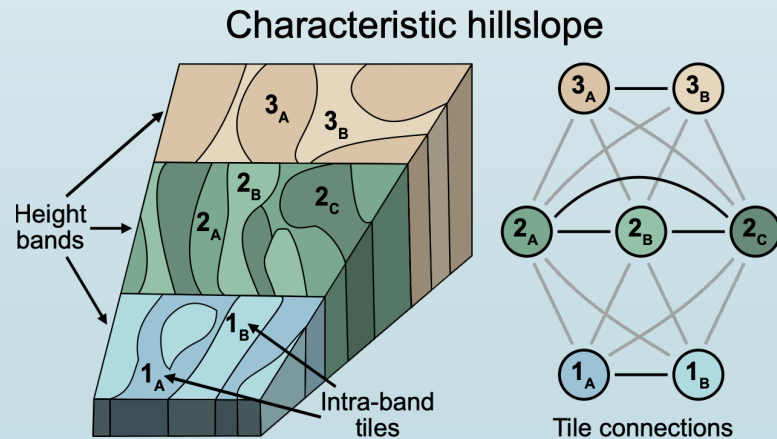
There is a long list of questions of how the complexity of land surface models enhances the initialization. But let's take a step back and ask, are these increases in complexity actually improving the modeling of surface fluxes and L-A interactions?

Role of Subsurface Flow on Surface States and Fluxes

4

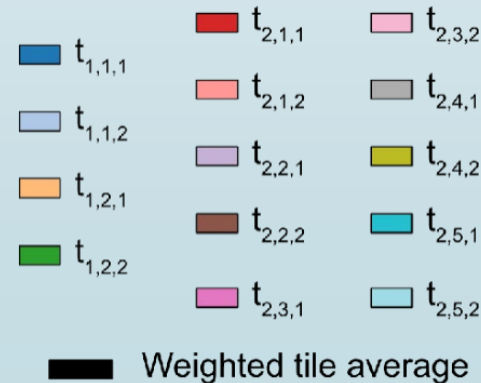
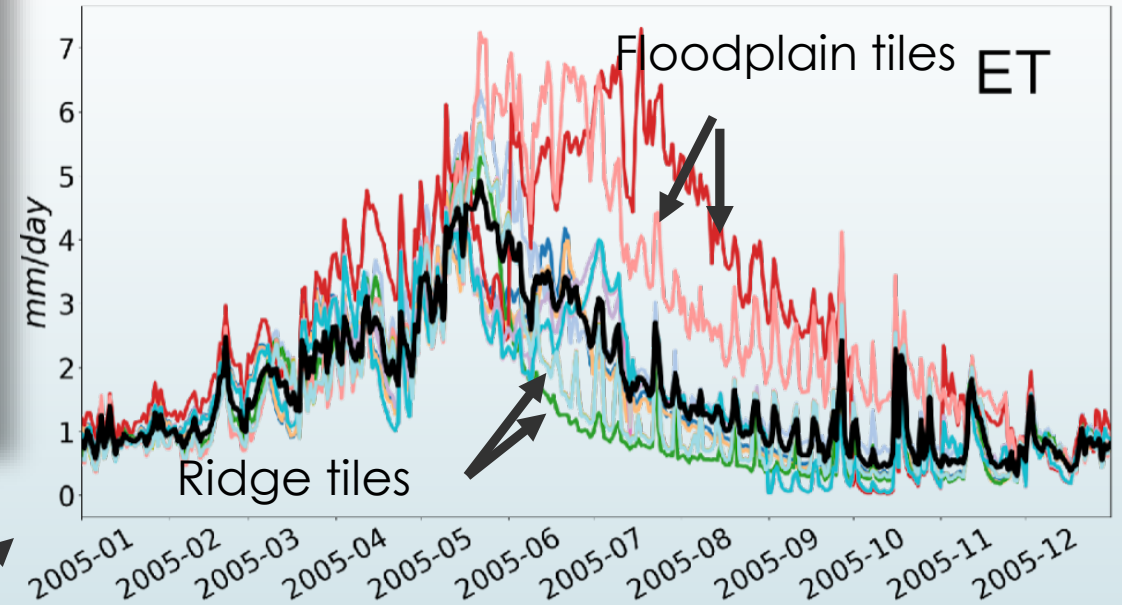


Credit: Stefan Kollet and Laura Condon



Chaney et al., 2018

Output from land model tiles of a 0.25 degree grid



Default Soil Maps
(Soil textures,
classes, etc.)

Model-specific
PTFs, or look-up
tables

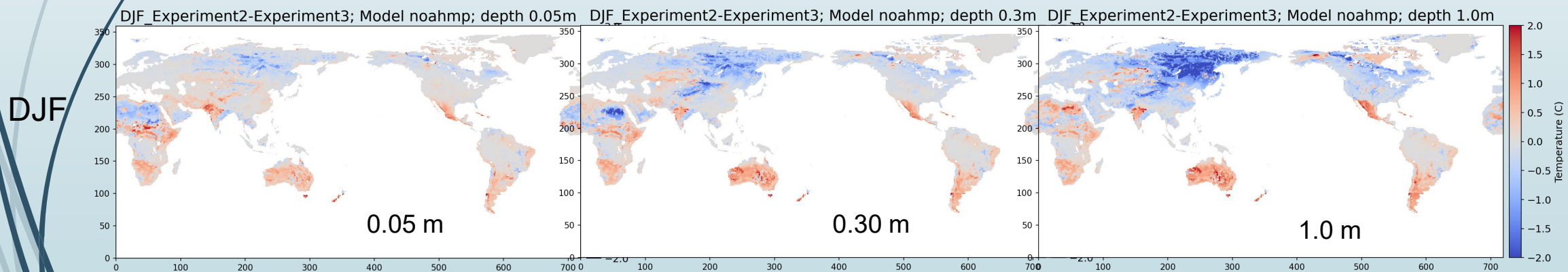
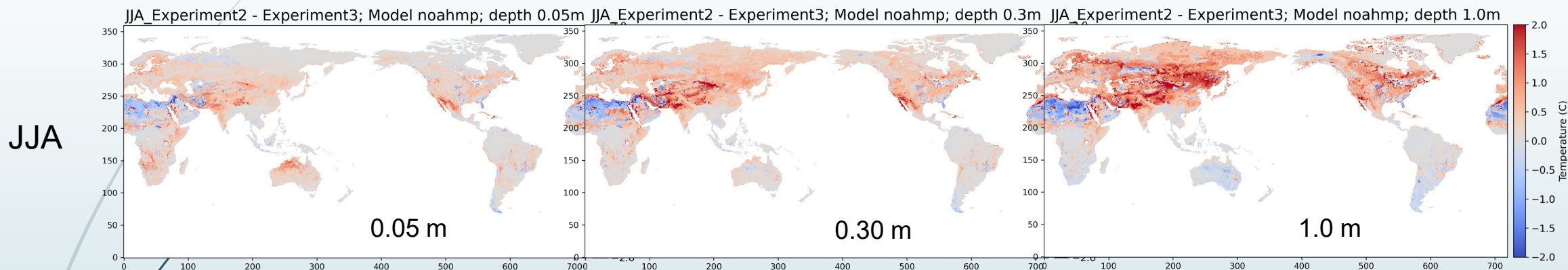
Soil Hydro-thermal
parameters for LSMs

5

SoilGrids Data
(Soil textures,
classes, etc.)

Role of Different Soil Maps

Noah-MP ($T_{EXP2} - T_{EXP3}$); 30-year daily average *difference*

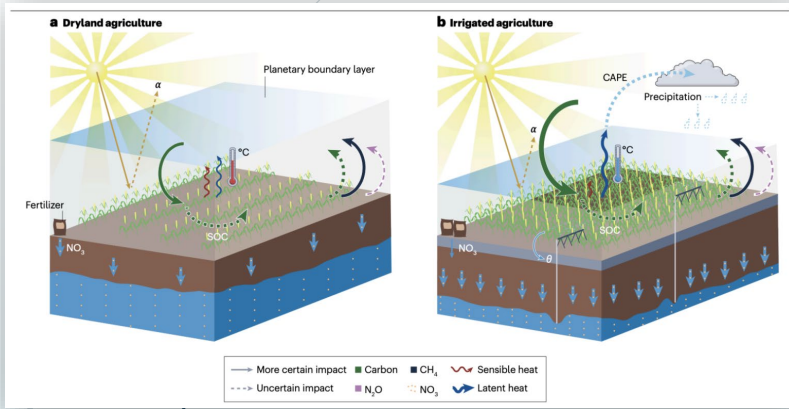


Increasing depth

Importance of Modeling Irrigation in Land-Atmosphere Interactions

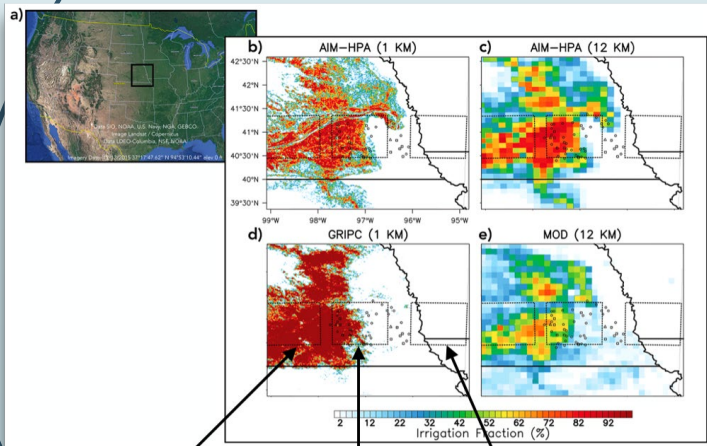
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Role of irrigation in surface energy balance

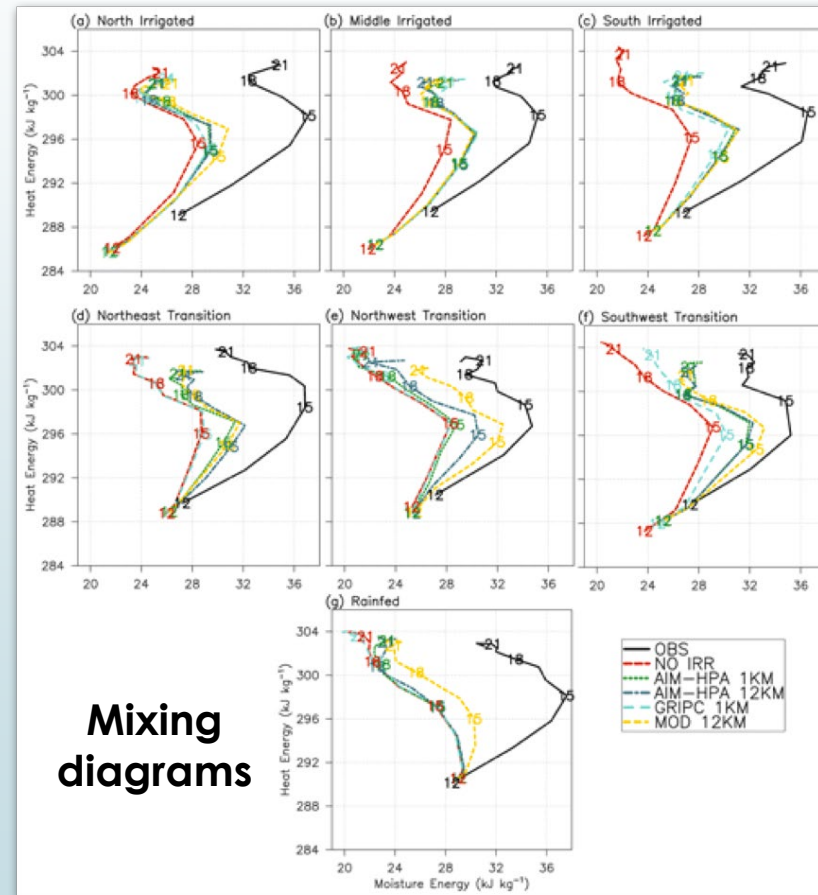


Sensitivity of macroscale response of atmosphere to different irrigation parameterizations/datasets

Case-study Noah-MP forced with different irrigation datasets



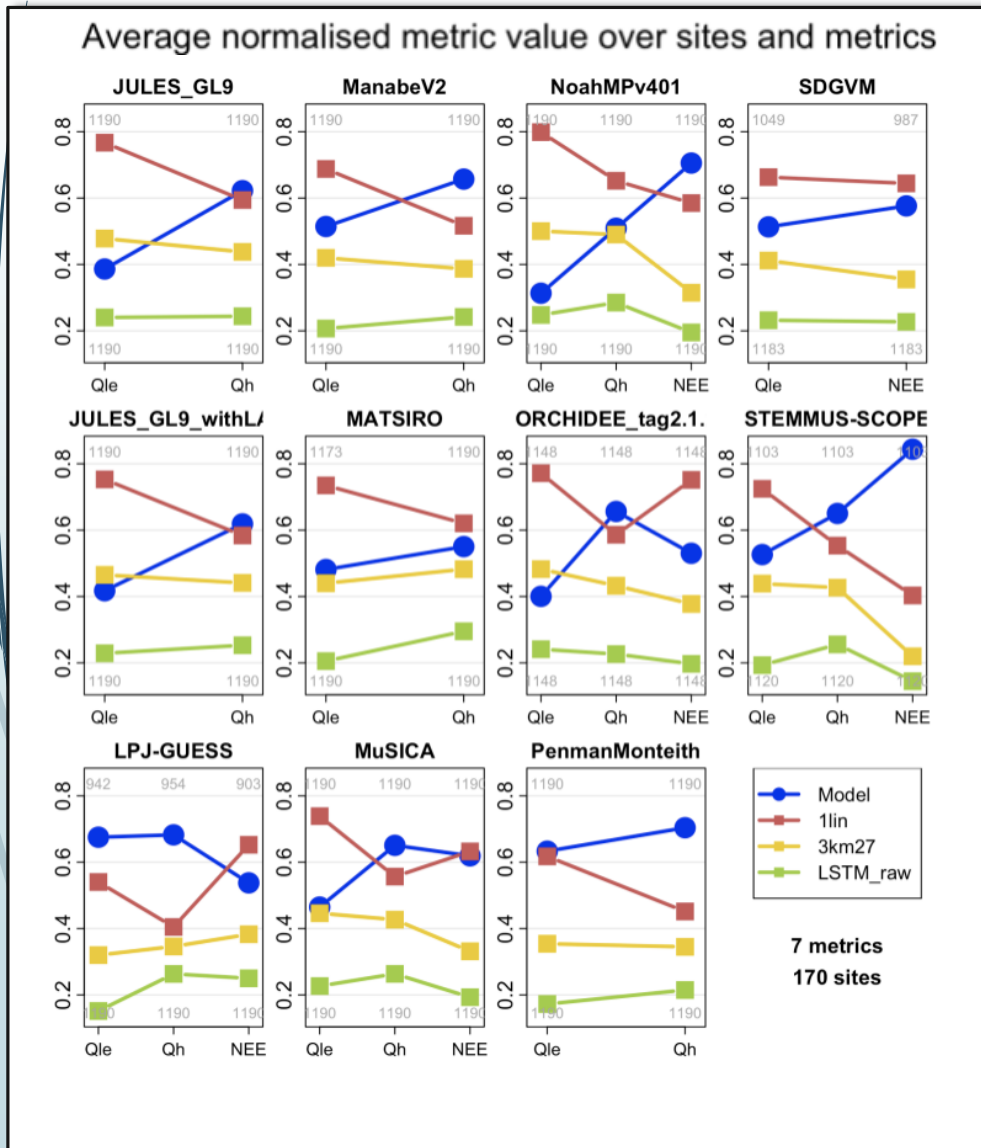
Irrigated Transition Rained



Credit: Tricia Lawston Parker

PLUMBER2: Benchmarking Surface Fluxes in Land Surface Models

7



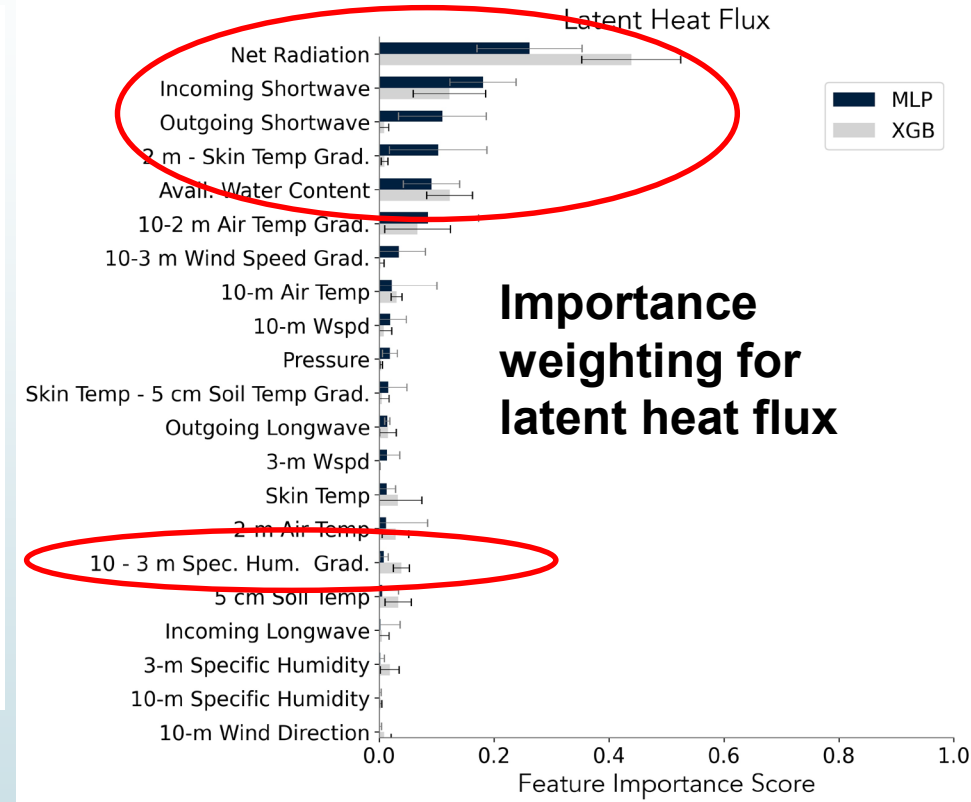
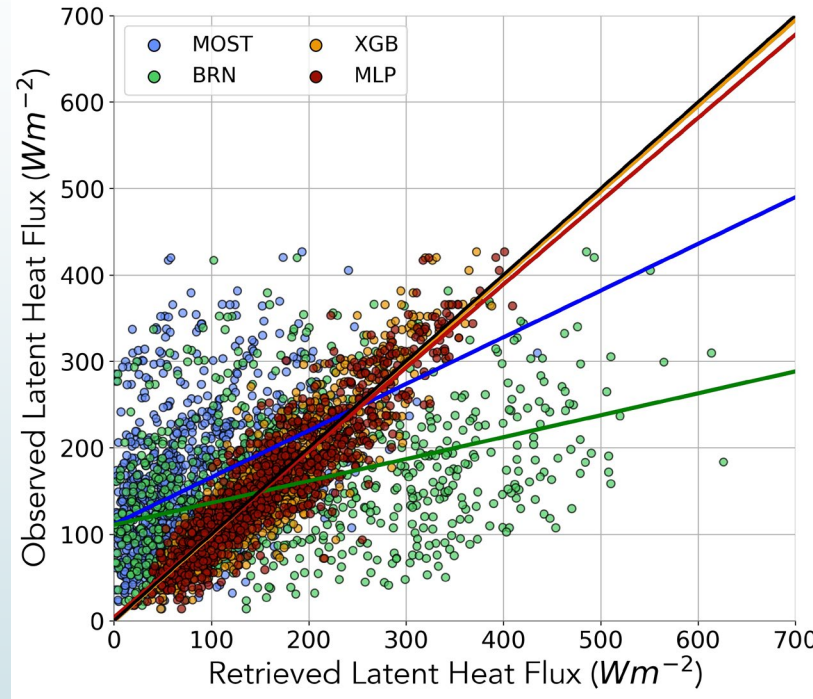
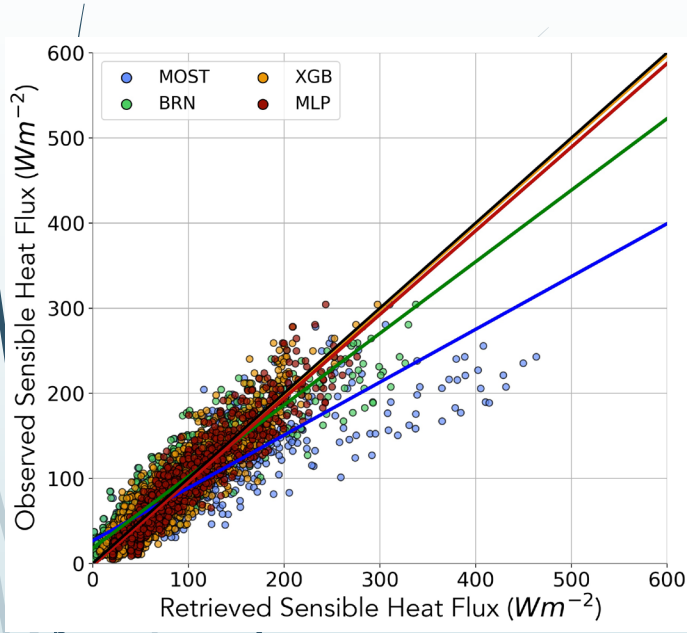
- Land surface models are run at 170 eddy covariance towers and the simulated fluxes are then compared to the observations.
- Q_e = Evaporation; Q_h = Sensible heat flux; NEE = Net ecosystem exchange
- Empirical approach (long-short term memory neural network; LSTM) provides the current “best case model” (i.e., benchmark). Uses same input data as land surface models.
- **Take home message:** Out-of-sample empirical approaches still appreciably outperform state-of-the-art land surface models

Credit: Gab Abramowitz

The Role of Monin-Obukhov Similarity Theory (MOST)

8

MOST is fundamental for the coupling of land surface and atmospheric model components. **It is applied in almost all coupled model systems from LES to global.**



$$\frac{\partial \bar{u}}{\partial z} = \frac{u^*}{\kappa z} \phi_M \left(\frac{z}{L} \right)$$

$$\frac{\partial \bar{\theta}}{\partial z} = \frac{H_0}{u^* \kappa z} \phi_H \left(\frac{z}{L} \right)$$

$$\frac{\partial \bar{q}}{\partial z} = \frac{Q_0}{u^* \kappa z} \phi_L \left(\frac{z}{L} \right)$$

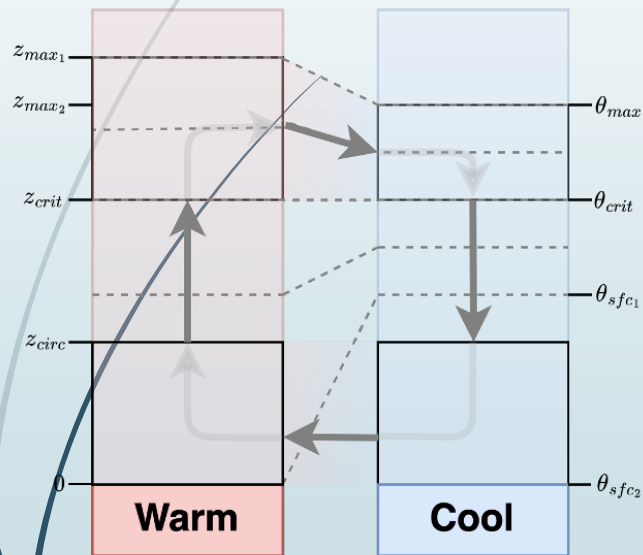
ML outperforms MOST and BRN. ML has great potential to improve surface layer flux relationships. Current partitioning of fluxes is huge error source (Lee and Buban JAMC 2020, Lee et al. MWR 2021, Lee and Meyers 2023, Wulfmeyer et al. BLM 2023).

Parameterizing Atmospheric Turbulence (Aerodynamic Resistances)

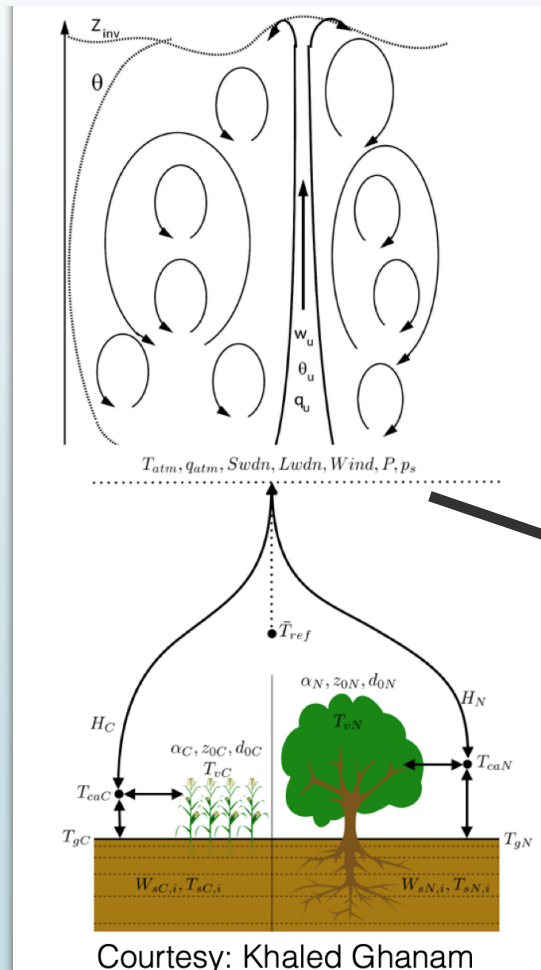
9

Surface heterogeneity-informed buoyant plumes

Parameterizing sub-grid secondary circulations



Courtesy: Tyler Waterman

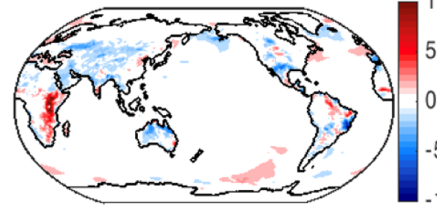


Courtesy: Khaled Ghanam

- Strong assumptions in surface turbulence parameterizations (e.g., planar homogeneity in MOST) that substantially limit their applicability.
- The **CLASP project** is exploring how surface coupling can be improved by accounting for parameterized heterogeneous-surface driven microscale to mesoscale circulations.

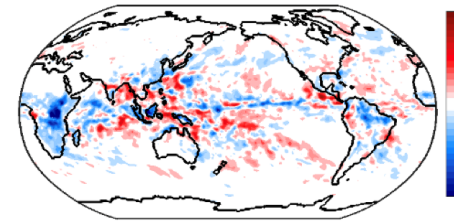
Global Climatology Differences (CLASP vs. Baseline)

Sensible Heat Flux (W/m²)
AM4-EDMF-HET minus AM4-EDMF
(MEAN = -0.08)

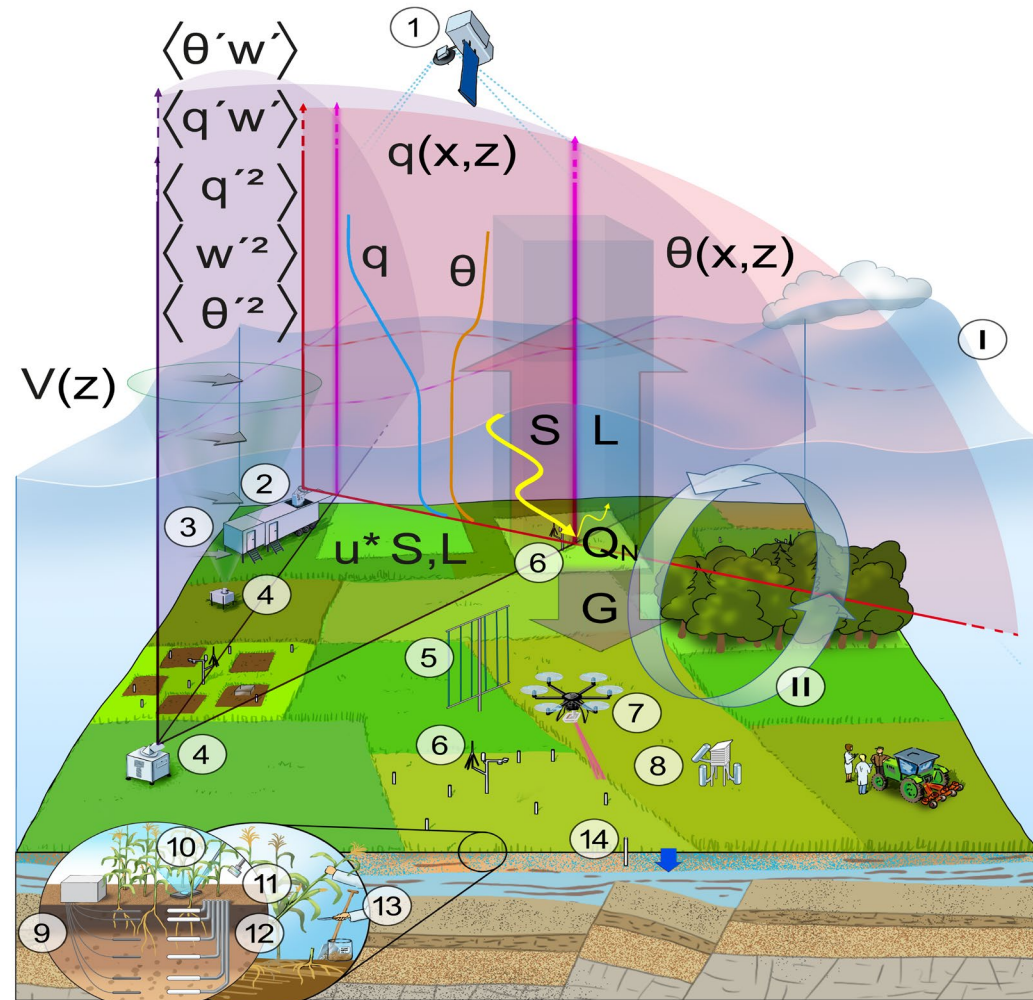


Precipitation (mm/day)

(f) AM4-EDMF-HET minus AM4-EDMF (DIFF = -0.01)



GEWEX LAFO (GLAFO) Design



Proposed sensor synergy:

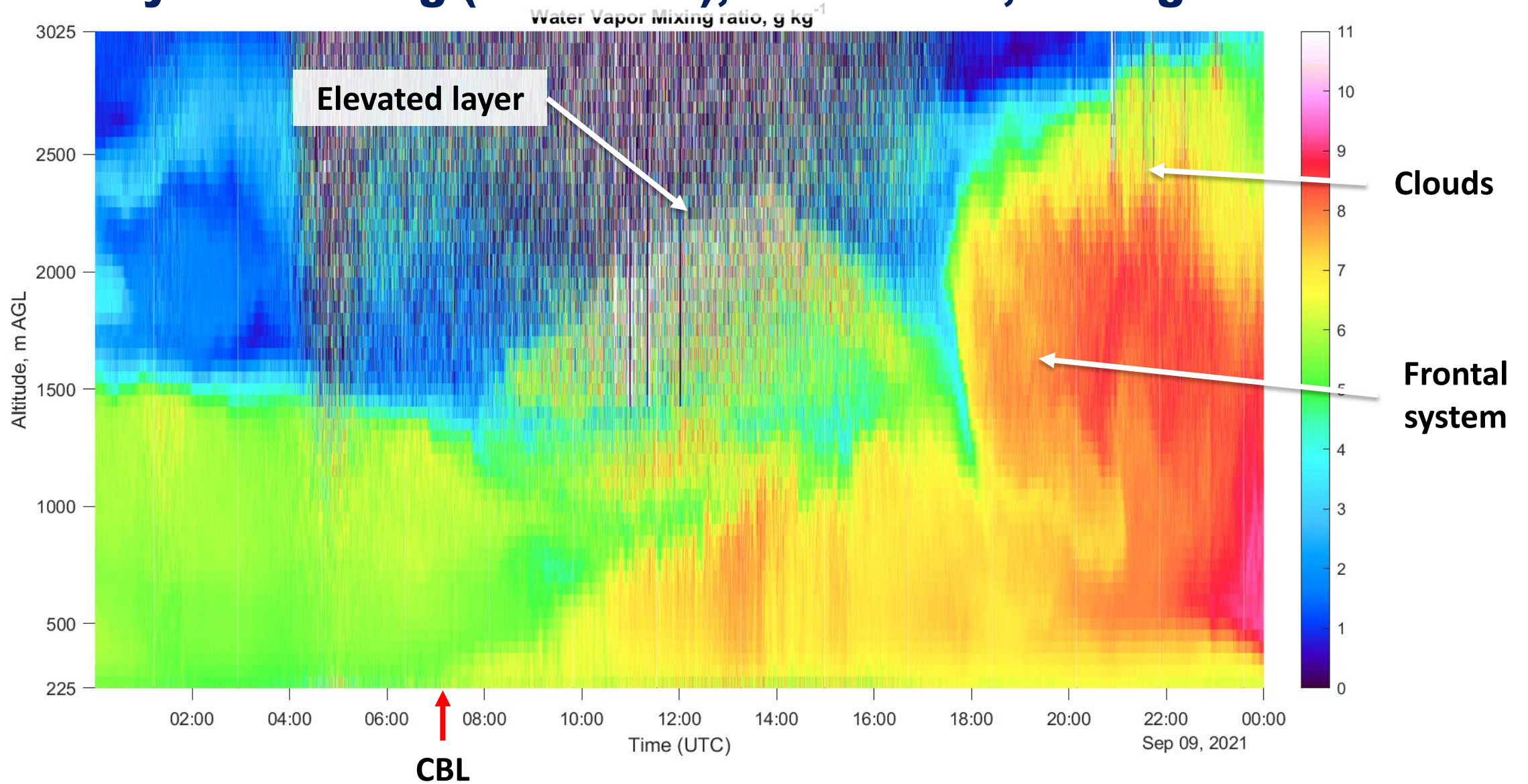
I: PBL top, II: mesoscale vortex

- 1: Satellite remote sensing
- 2: Vertically staring Doppler, water vapor, temperature, and CO₂ lidar systems, infrared spectrometer (IRS), microwave radiometer (MWR), cloud radar
- 3: Scanning Doppler, water vapor, temperature, and CO₂ lidar systems
- 4: Scanning Doppler lidar systems
- 5: Fiber-based distributed sensors
- 6: Energy balance and eddy covariance stations
- 7: Unmanned aerial vehicle (UAV)
- 8: Water vapor and CO₂ isotope sensor
- 9: Time-domain reflectometers (TDRs)
- 10: Leaf area index (LAI) measurement
- 11: Gas exchange system for photosynthesis and transpiration rates
- 12: Tensiometers
- 13: In-situ canopy measurements such as biomass and canopy height
- 14: Soil moisture and temperature network

Wulfmeyer et al. GEWEX Newsletter 2020, GLAFO White Paper 2021 (see <https://www.gewex.org/panels/global-landatmosphere-system-study-panel/glass-projects>)

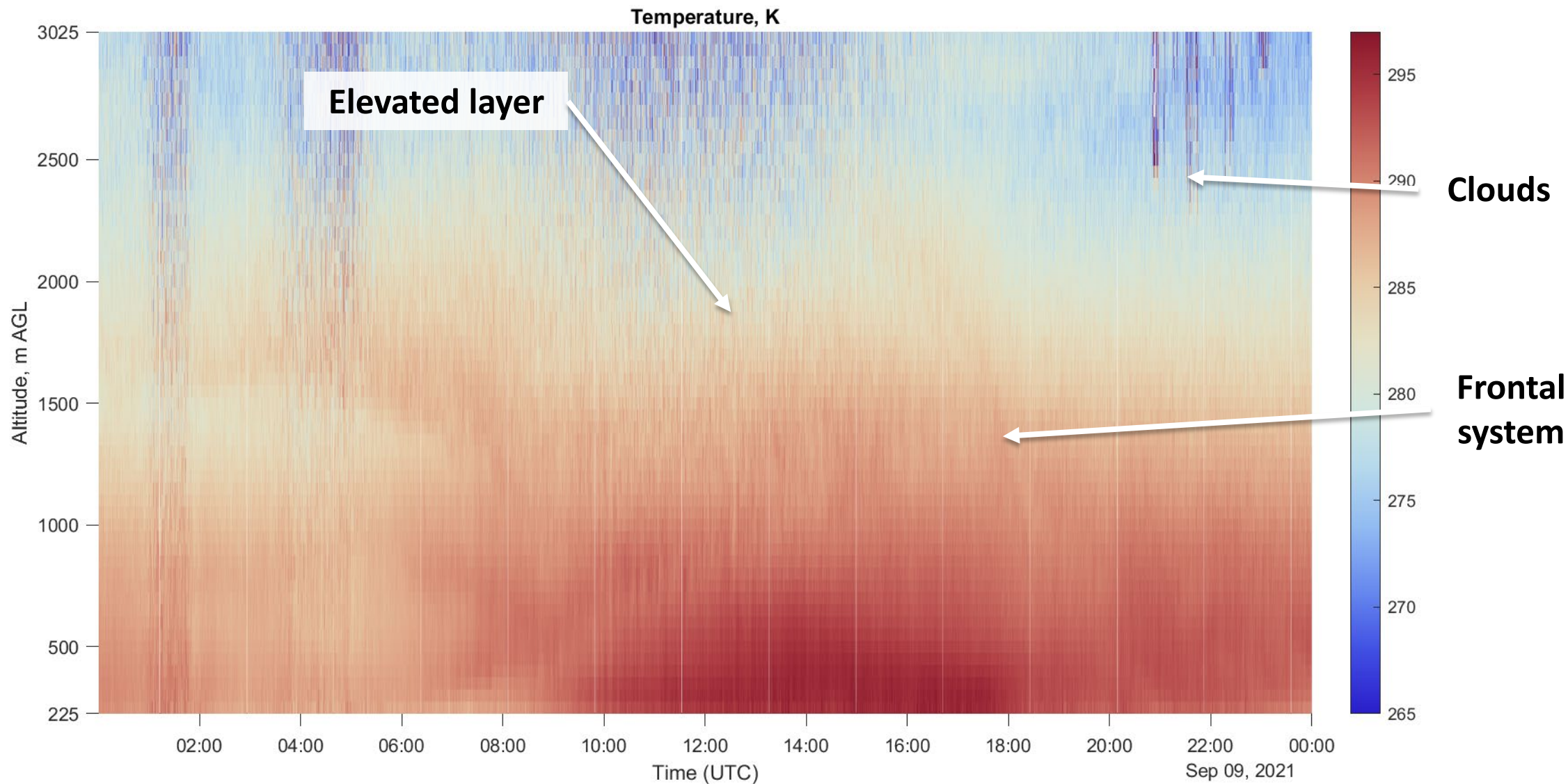
German Weather Service (DWD) Meteorological Observatory Lindenberg (MOL-RAO), 09.09.2021, Mixing Ratio

10 s, 50 m



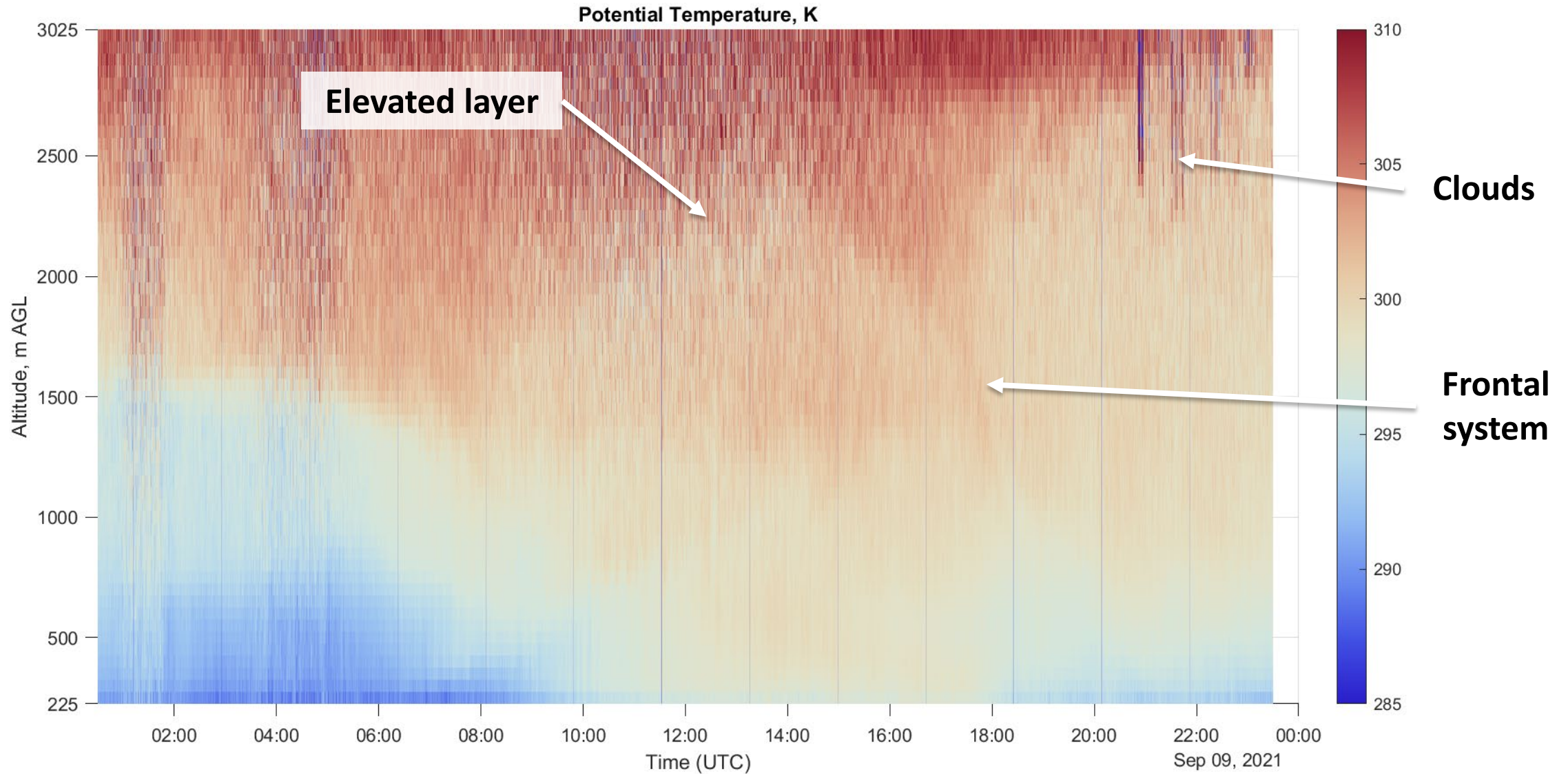
Temperature

10 s, 50 m



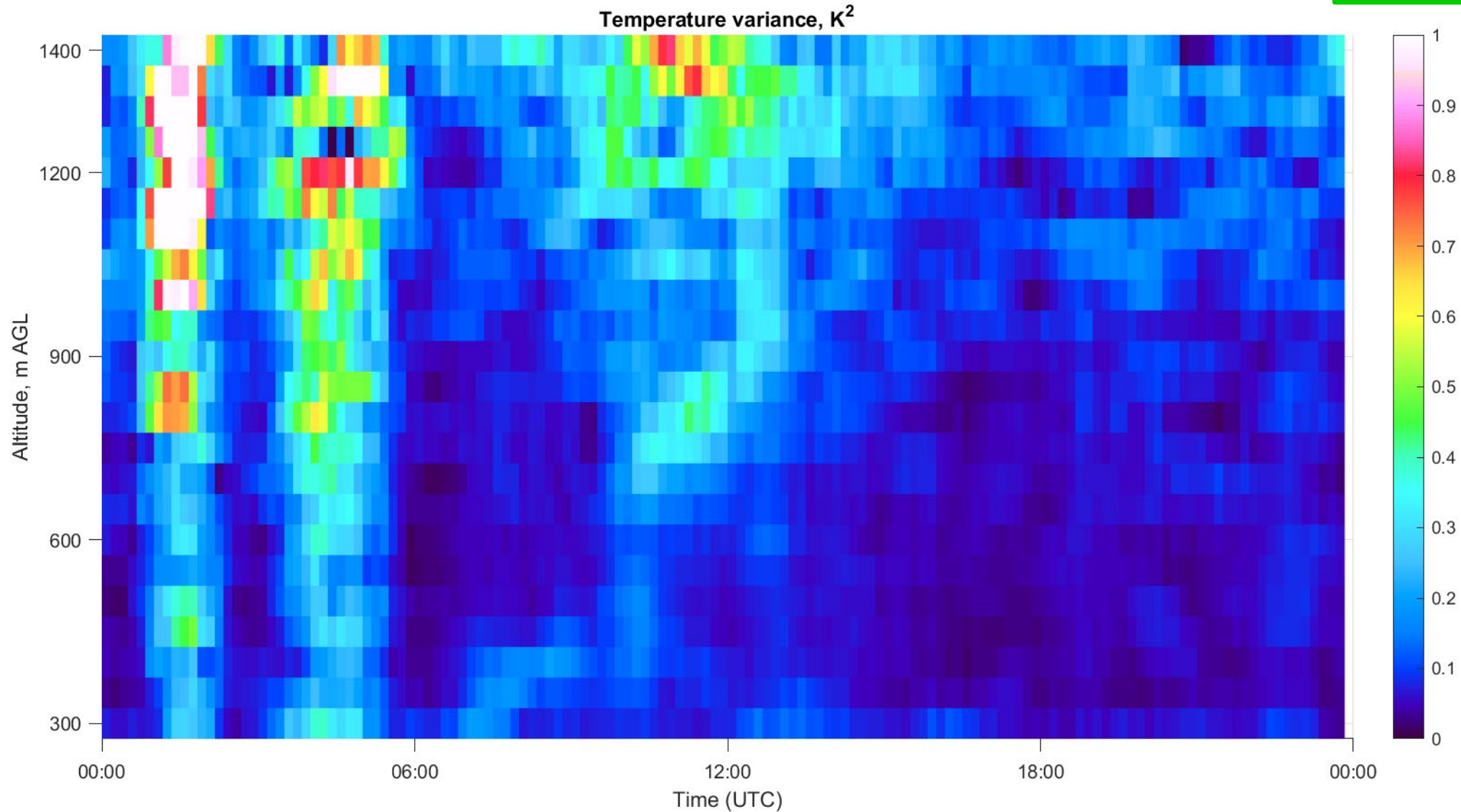
Potential Temperature

10 s, 50 m



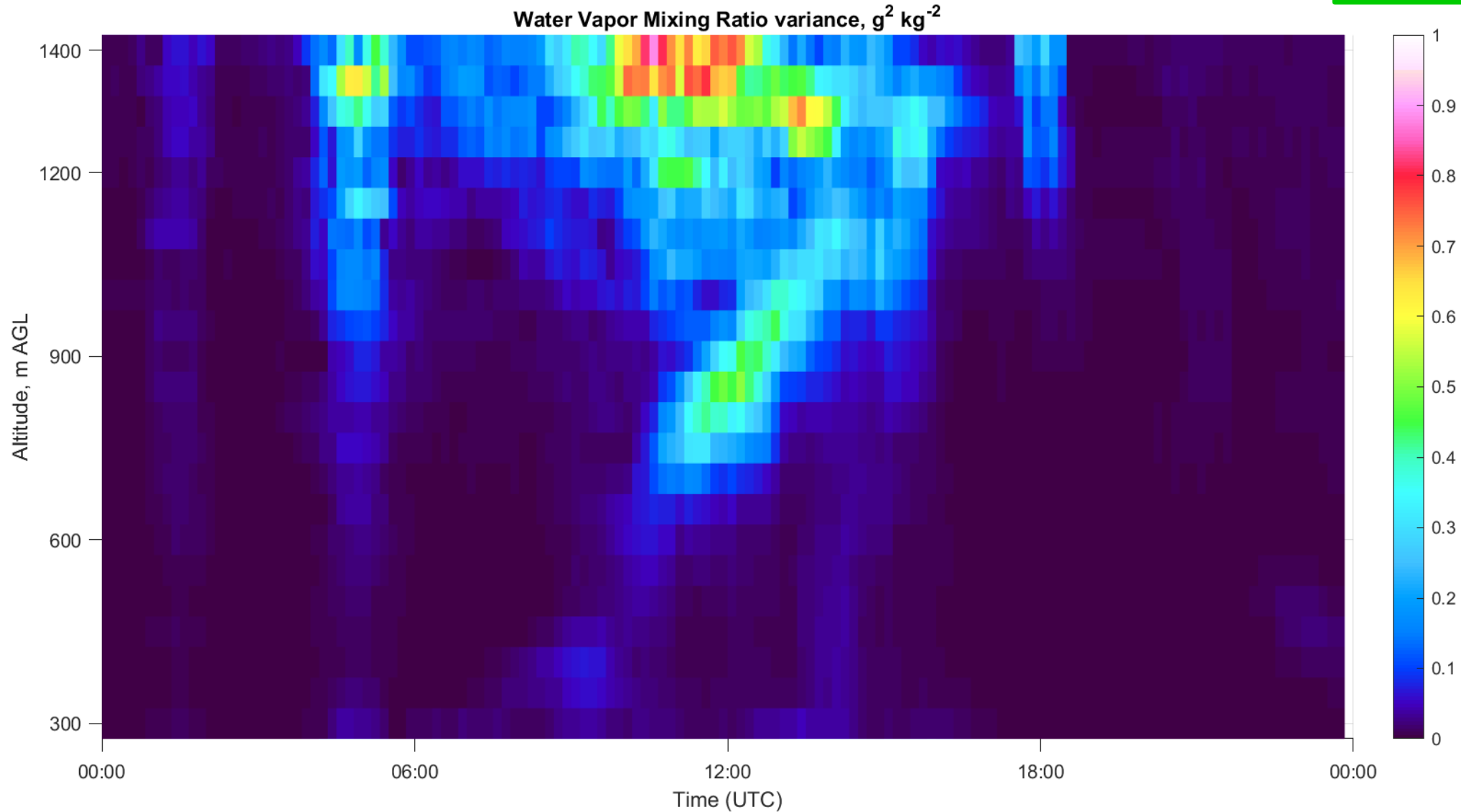
Temperature Variance

1 h (averaged every 600s), 50 m



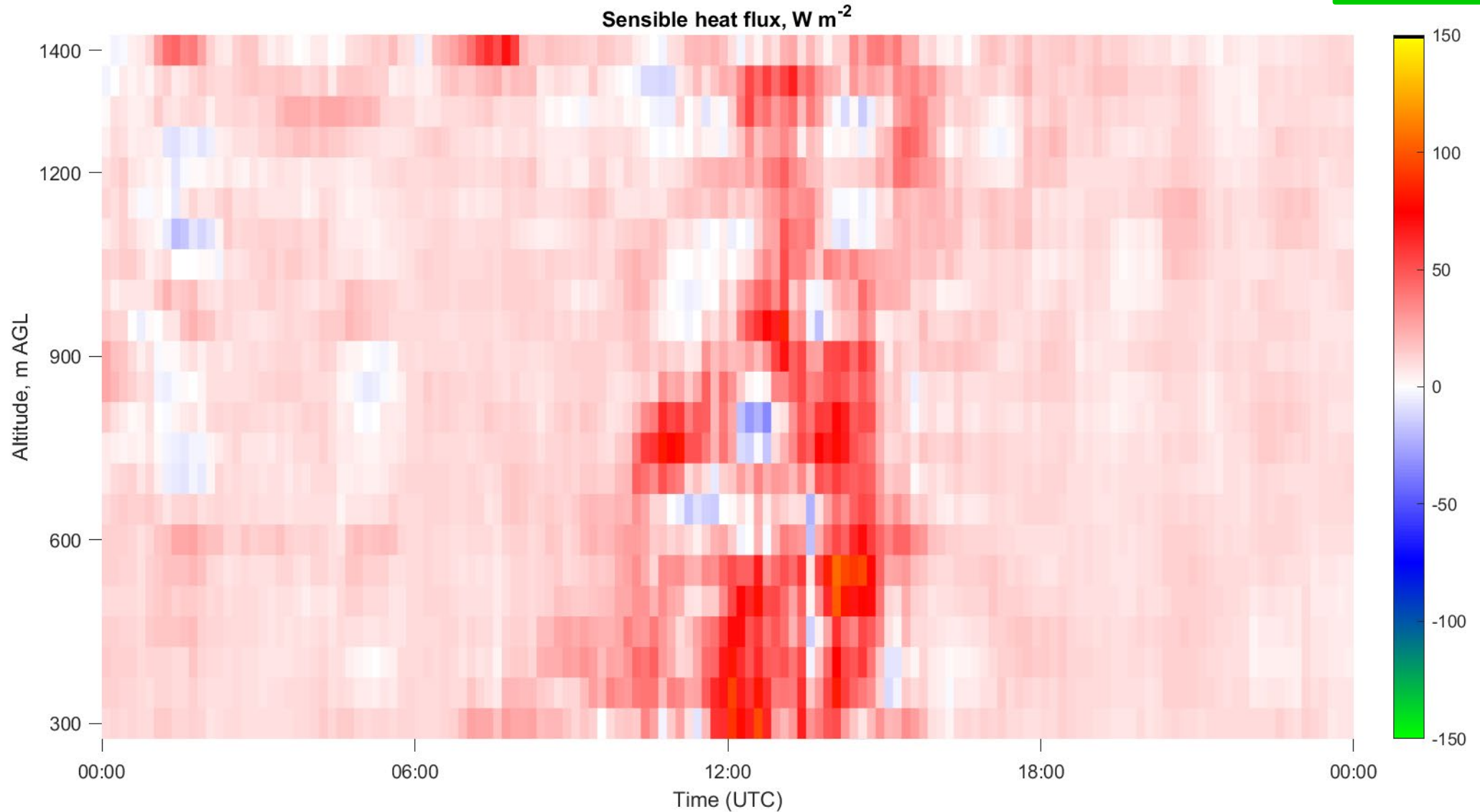
Mixing Ratio Variance

1 h (averaged every 600s), 50 m



Sensible Heat Flux

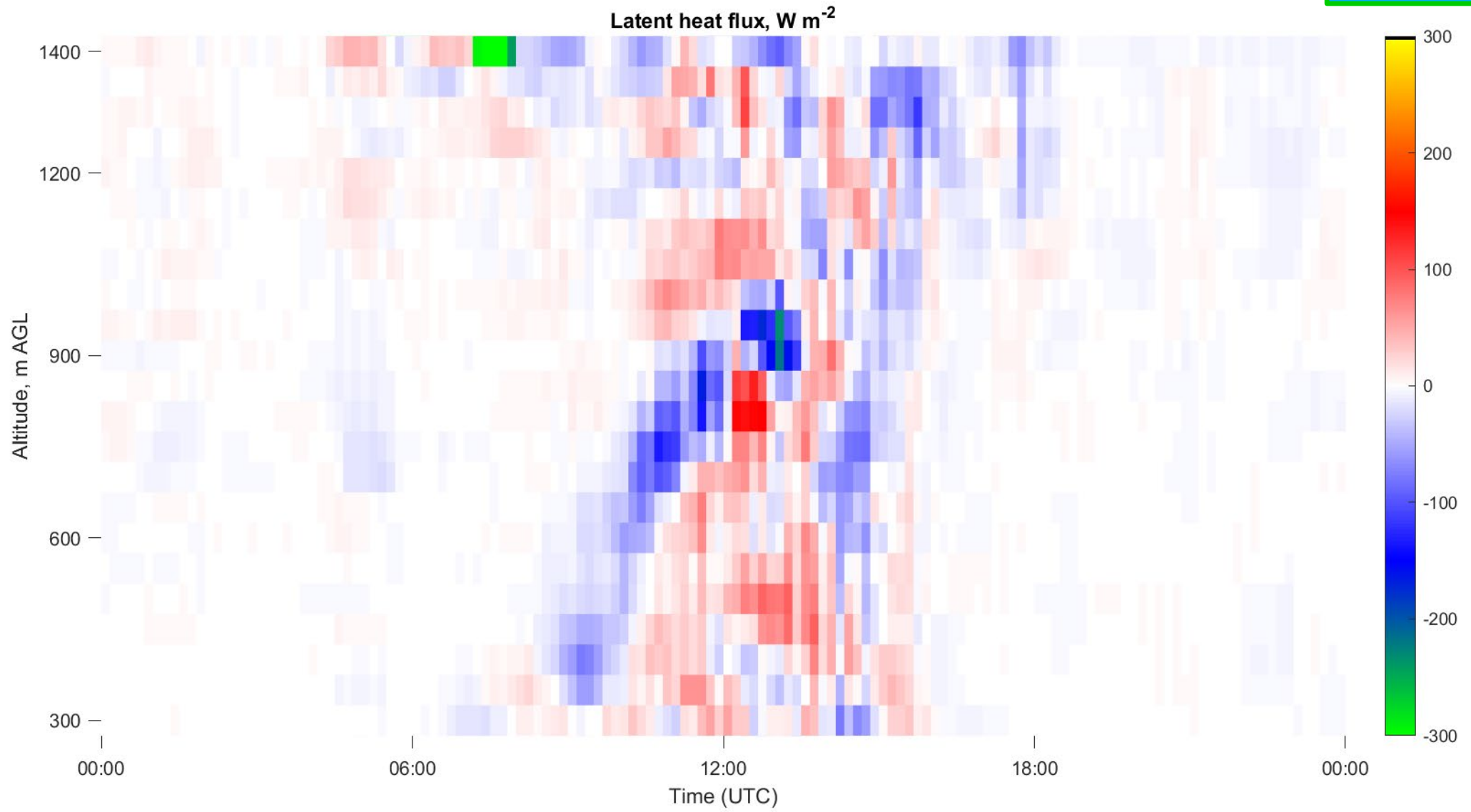
1 h (averaged every 600s), 50 m



Latent Heat Flux

UNIVERSITÄT HOHENHEIM

1 h (averaged every 600s), 50 m

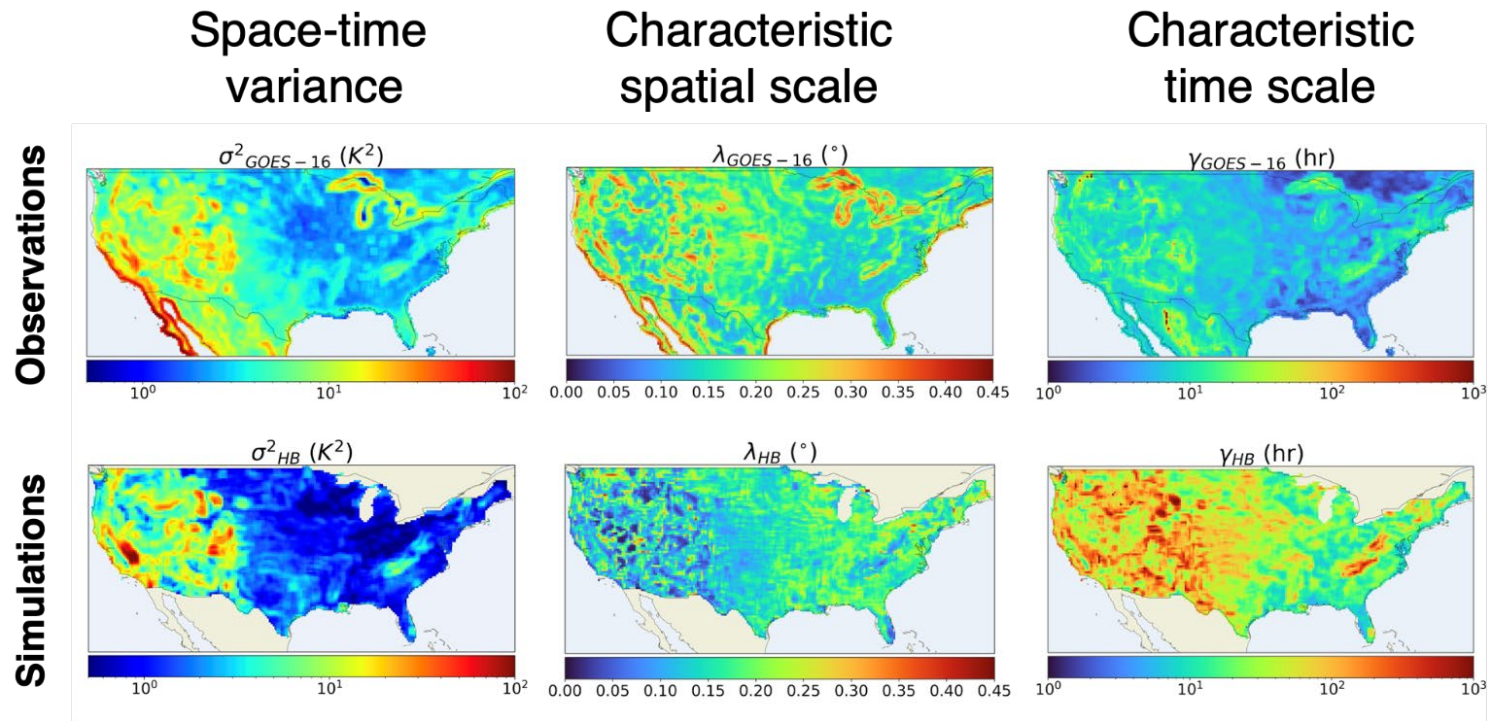


Need to Move to Evaluate Space-Time Patterns of Land Surface Models

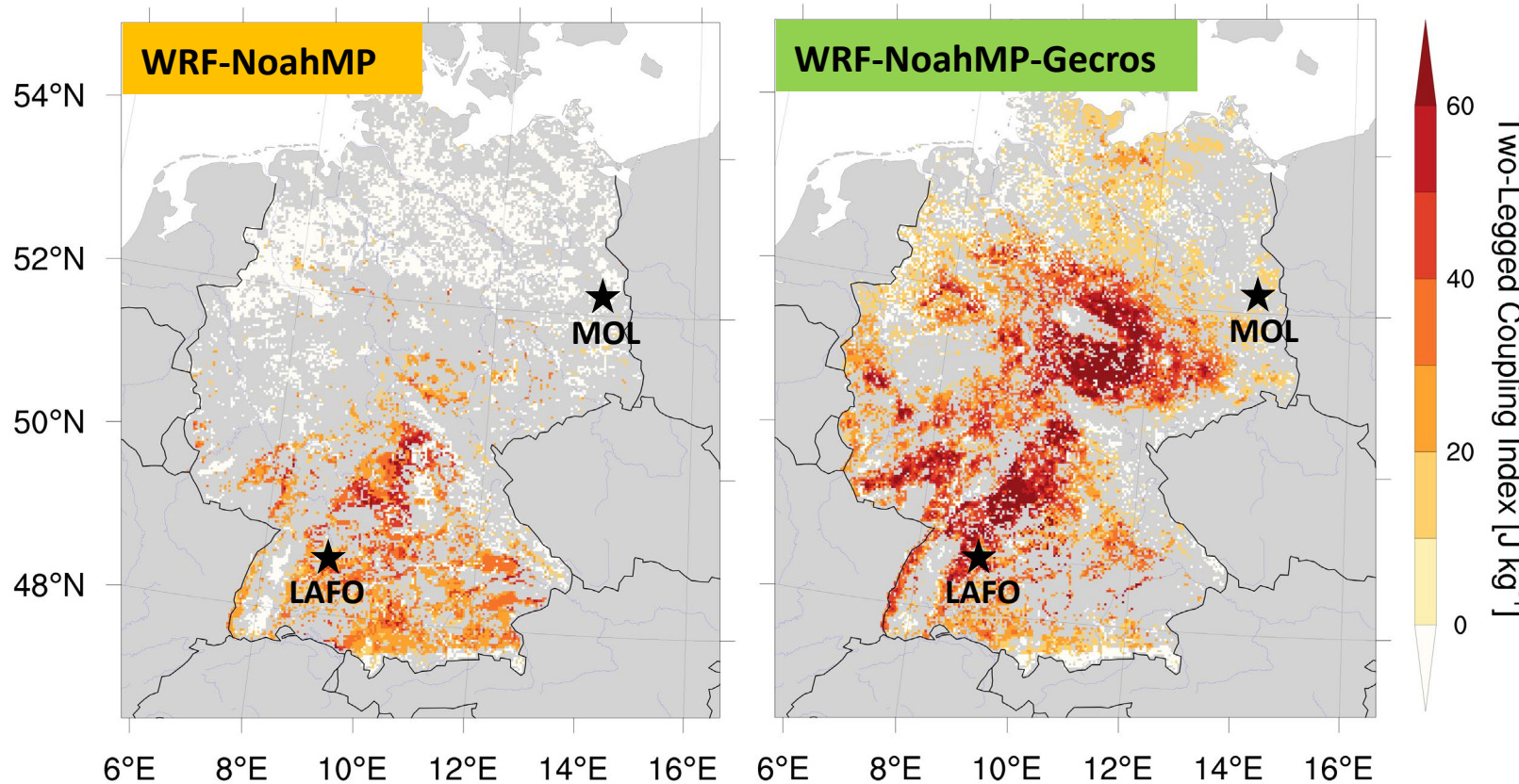
18

- Land surface models produce "movies" of states and fluxes. These movies should be evaluated as such, instead of only in space or time
- **CLASP** is exploring how space-time covariance on 10 min – hourly remotely sensed global Land Surface temperature (LST) data can be used to explore the realism of simulated space-time patterns of evaporation.

$$C(h, \tau) = \sigma^2 e^{-\left(\frac{\tau}{\gamma}\right)^a - \left(\frac{h}{\lambda}\right)^a}$$



How do State-of-the-Art Models Represent Land-Atmosphere (L-A) Feedback?



April-August 2005, grid increment 3 km (Warrach-Sagi et al. JGR 2022)

$$TLCI = \sigma(q_{l,s}) \frac{d ET}{dq_{l,s}} \frac{d CAPE}{d ET}$$

TLCI: Two legged-coupling index

σ : Standard deviation

$q_{l,s}$: Soil moisture content

ET: Evapotranspiration

CAPE: Convective available potential energy

White: No two-legged coupling

Grey: Not analyzed due to lack of observations outside Germany or no significance

- Current L-A system models seem to underestimate the sensitivity of cloud formation to land-surface properties.
- This requires an understanding of land-atmosphere (L-A) feedback on the km-scale.

Summary and Outlook

- Accurate representation of **L-A feedback** crucial for model initialization.
- Model systems should be verified with respect to **surface fluxes** and various **feedback metrics**.
- **Correct soil maps** critical for accurate simulations of soil temperatures.
- **Irrigation** needs to be considered and simulated in respective regions.
- **Increase of complexity of LSM does not improve the simulation of fluxes**. The best approach is a simple ML effort.
- **MOST introduces large errors in the simulation of surface latent heat fluxes**, which translates to incorrect flux partitioning. **The simulation of surface fluxes must be replaced**.
- **Dynamics of vegetation must be considered for simulations on the seasonal scales**.
- **CLASP** developed **new parameterization of turbulence** over complex terrain.
- **GLAFOs** provide **turbulence and flux profiles** for process studies and model verification.

21

Thank you